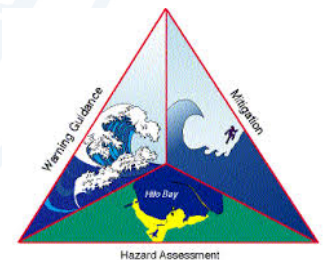




UNIVERSITY *of* DELAWARE

Landslide Tsunami Benchmark Workshop

Intercomparison of Model Results





Categorizing the Models

- Models characterized by approach to hydrodynamics:
 - ○ Hydrostatic (depth-integrated or not)
 - ☆ Weakly dispersive (Boussinesq-type)
 - ⚡ Fully dispersive (non-hydrostatic or mild slope)
 - * Navier-Stokes solvers
- Models characterized by approach to modeling slide motion:
 - ○ Solid slide or analytic specification of bottom as $f(t)$.
 - ○ Depth-integrated slide, Newtonian viscous rheology
 - ○ Depth-integrated slide, granular rheology
 - ○ 3D slide, continuous density variation, variable rheology



Model name	Model #	Hydro Type	Slide Type	Numerical	BM1	BM2-d61	BM2-d120	BM3-case30	BM3-case32	BM4-test17	BM5 Case 1	BM5 Case 2	BM6	BM7
data	0	0	0		y	y	y	y	y	y	y	y	y	y
nhwave-s	1	3	1	2	y	y		y	y					
nhwave-v	2	3	2	2						y				y
nhwave-g	3	3	3	2						y	y	y		
globouss-l	4	2	1			y	y							
globous-nl	5	2	1			y	y							
boussclaw-0	6	2	1			y	y							
boussclaw-15	7	2	1											
tsunami3d	8	4	4			y	y			y				y
thetis-newt	9	4	4					y		y	y			
nhwave-3d	10	3	4	2						y				
thetis-mui	11	4	4								y			
ls3d	12	2	1			y	y							y
2lcmflow	13	1	3	2						y	y			
Alaska GI-L	14	1	2			y	y			y				y
nhwave-s-hydrostatic	15	1	1	2		y	y							
nhwave-g-hydrostatic	16	1	3	2						y				
landslide-hysea	17	1	3			y	y			y				y
fbslide	18	1	1											
Lynett-nlsw	19	1	1			y	y							
Lynett-Bouss	20	2	1			y	y							
Lynett-MSE	21	3	1			y								
NHWAVE+FUNWAVE	22	3	2											y

Hydro Types

- 0 - data
- 1 - non-dispersive
- 2 - weakly dispersive, one layer
- 3 - nonhydrostatic
- 4- Navier Stokes

Slide Types

- 1 - moving boundary
- 2 - viscous layer
- 3 - granular layer
- 4 - 3D multiphase



Analyzing Results: Benchmarks 2 and 4

- RMS error measures comparing models to data
- Behavior of transient solutions in frequency and time
- Distributions of maximum and minimum elevations, maximum excursions, and variance of generated wave packets

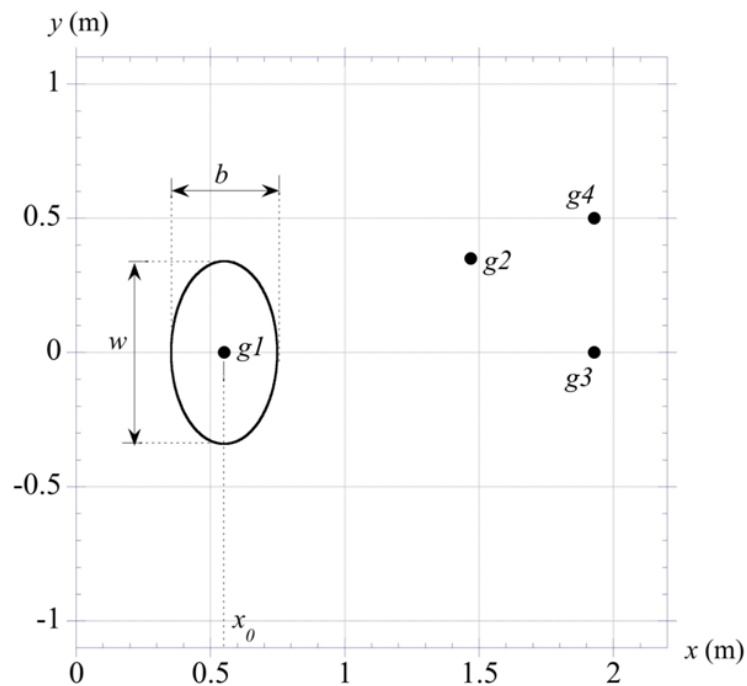


RMS error estimates

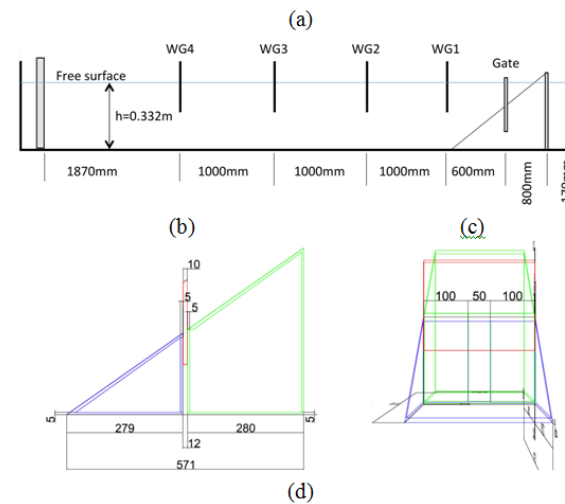
- For both benchmarks, model and data times are synchronized by aligning the maximum crest elevation occurring at Gauge 2.
 - For benchmark 2, hydrostatic cases were synchronized at minimum elevation at Gauge 1
- Two error estimates computed, one for each gauge alone and one over the entire set of gages. Only the first option is presented here.



Gauge positions for Benchmarks 2 and 4



Benchmark 2

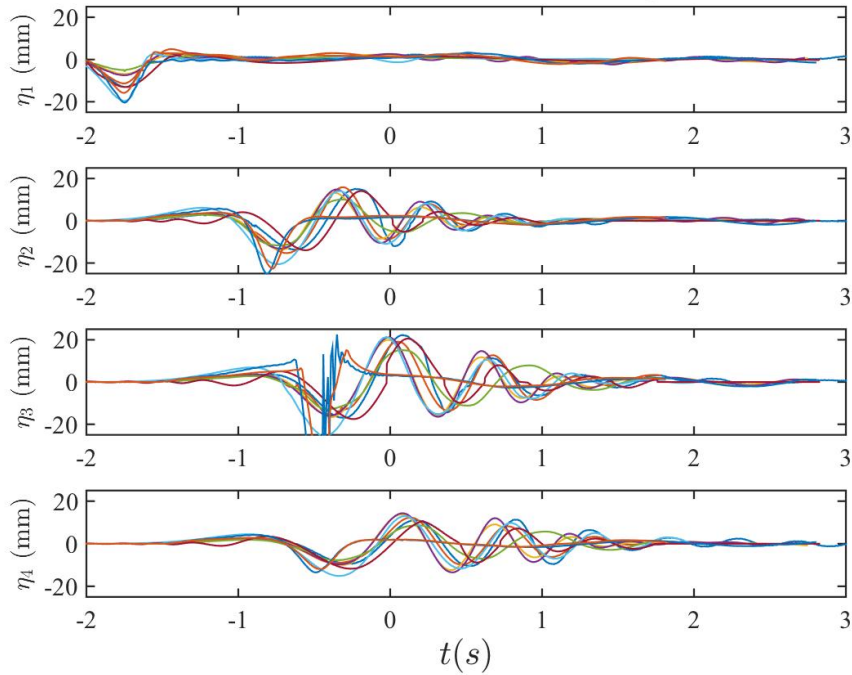


Benchmark 4

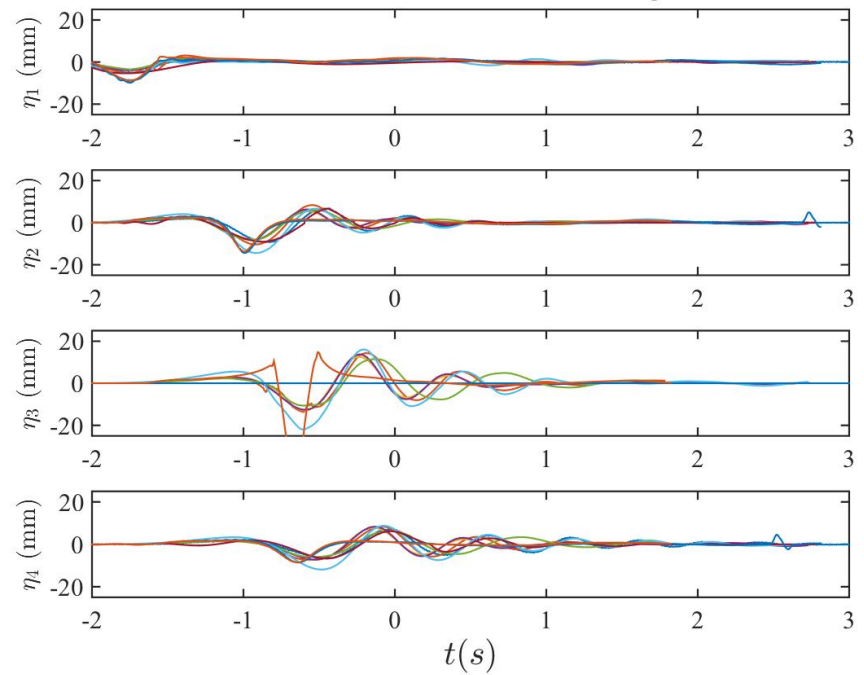


Benchmark 2: Overall scatter in results

BM 2: Models 0 1 4 5 6 8 12 14 15, Submergence $d = 061$

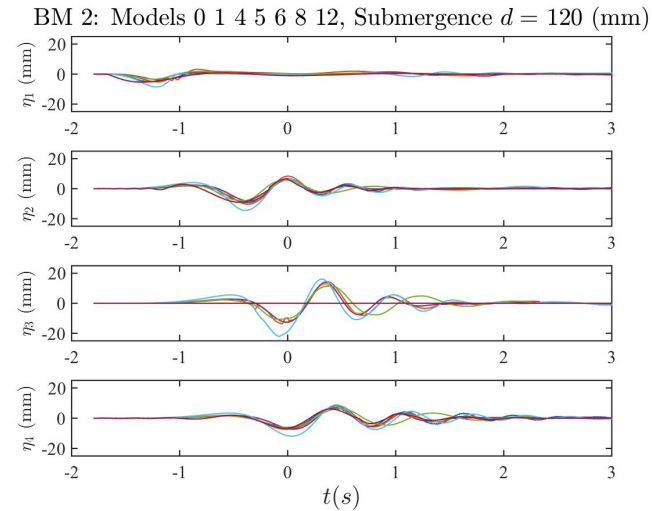
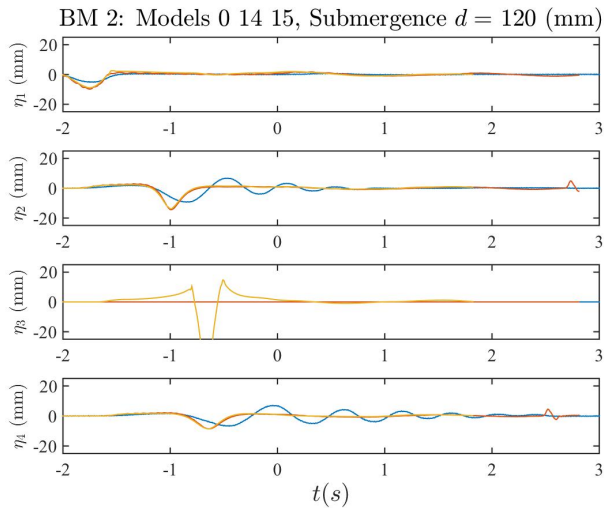
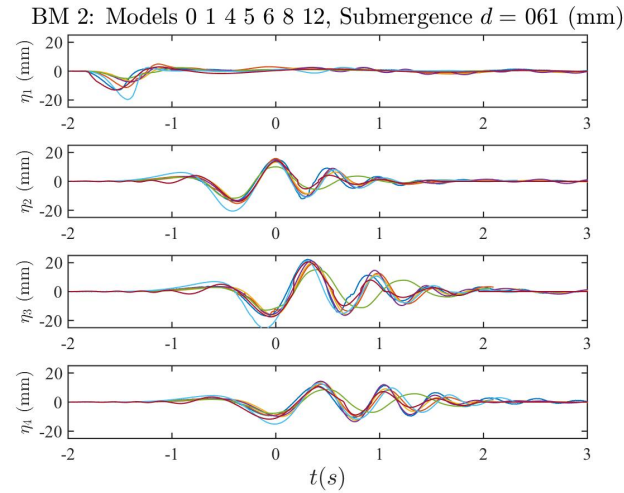
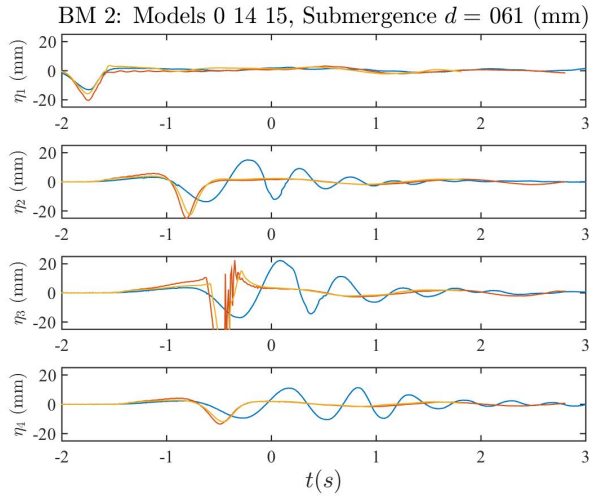


BM 2: Models 0 1 4 5 6 8 12 14 15, Submergence $d = 120$ (mm)





Benchmark 2: Hydrostatic vs. nonhydrostatic cases



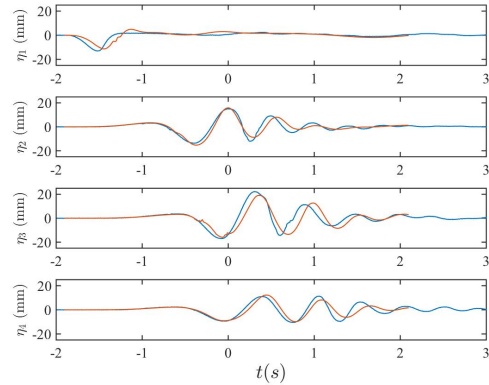
hydrostatic

Non-hydrostatic

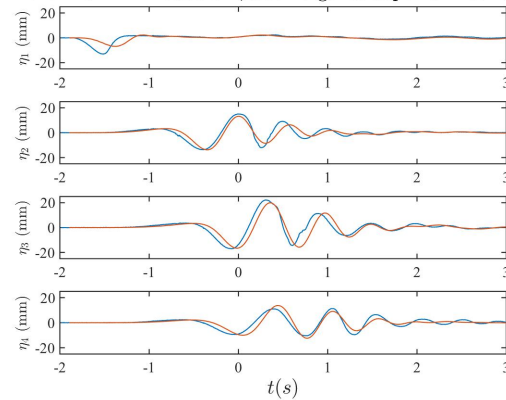


Benchmark 2: Individual results for dispersive models

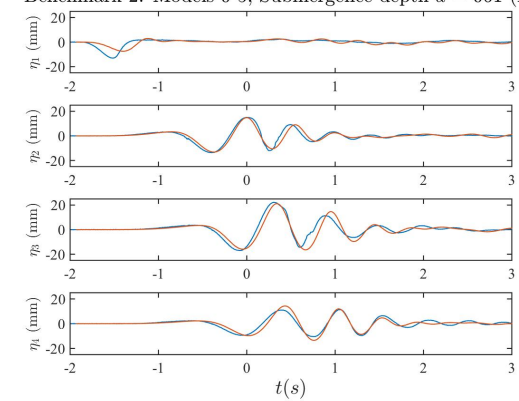
Benchmark 2: Models 0 1, Submergence depth $d = 061$ (mm)



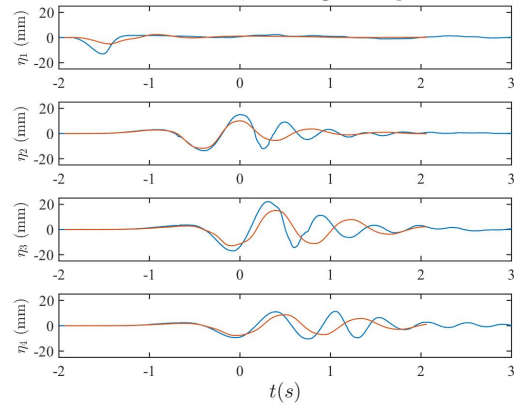
Benchmark 2: Models 0 4, Submergence depth $d = 061$ (mm)



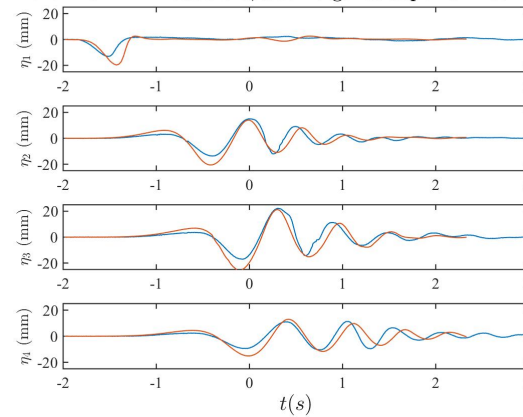
Benchmark 2: Models 0 5, Submergence depth $d = 061$ (mm)



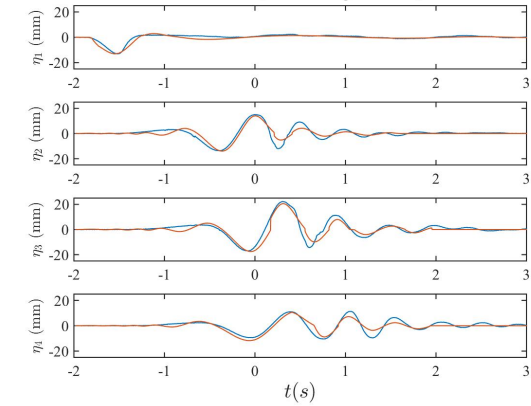
Benchmark 2: Models 0 6, Submergence depth $d = 061$ (mm)



Benchmark 2: Models 0 8, Submergence depth $d = 061$ (mm)

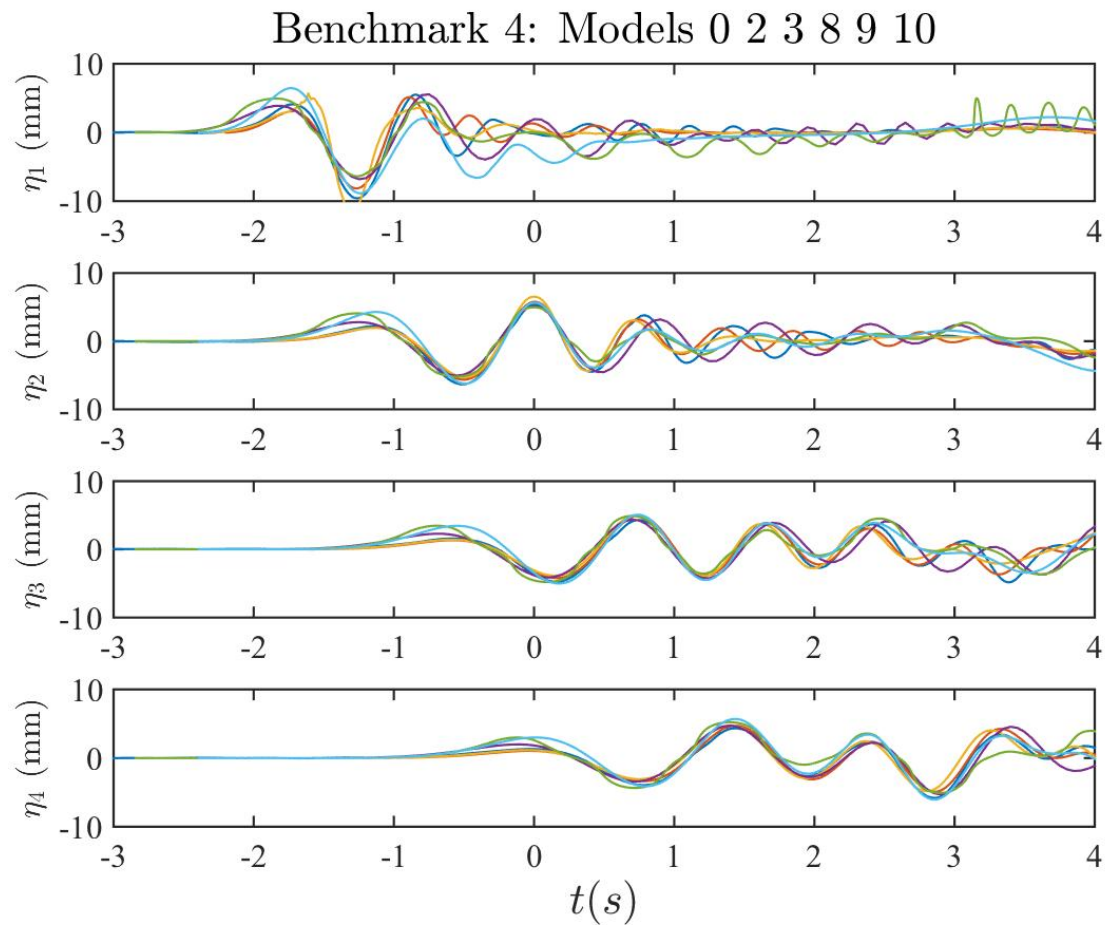


Benchmark 2: Models 0 12, Submergence depth $d = 061$ (mm)





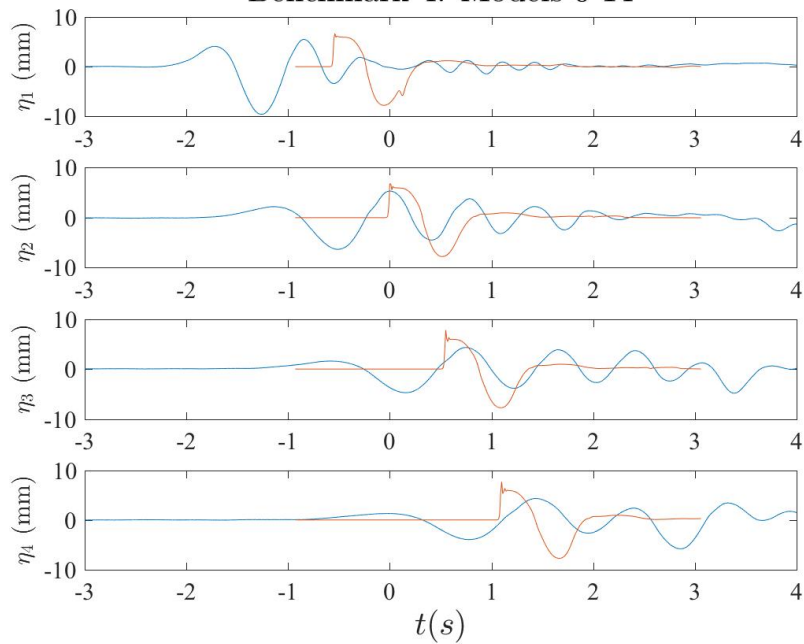
Benchmark 4: Dispersive models



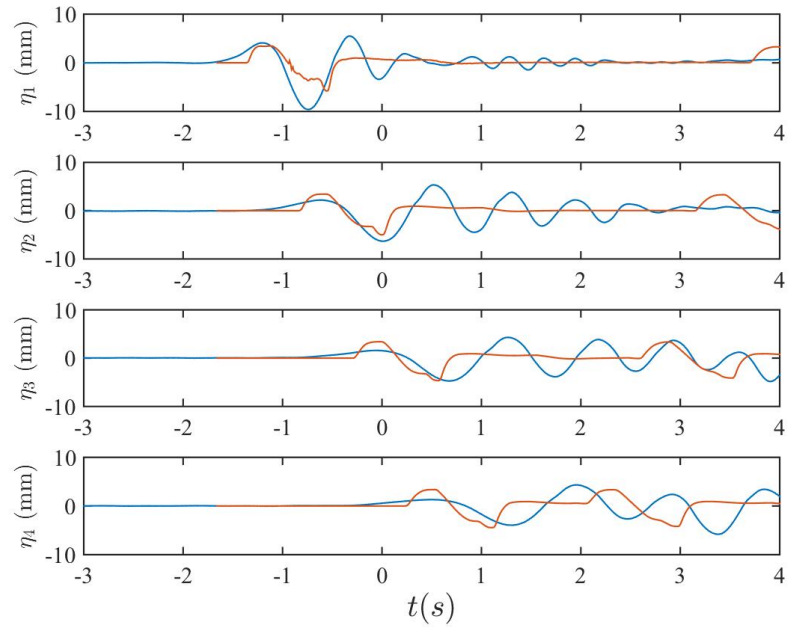


Sample nondispersive model results

Benchmark 4: Models 0 14

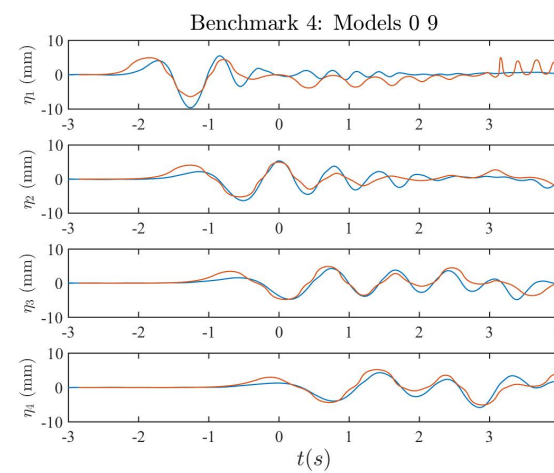
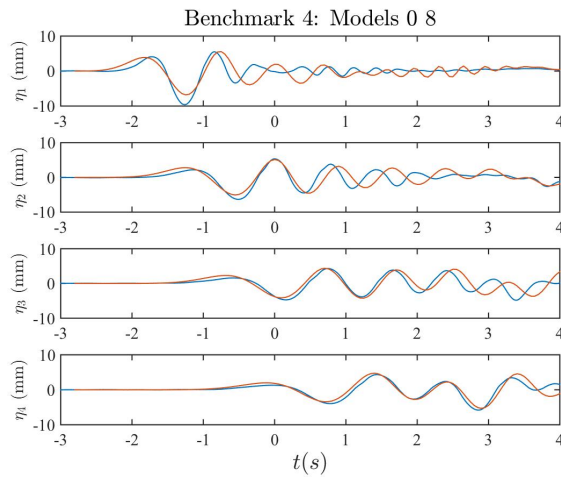
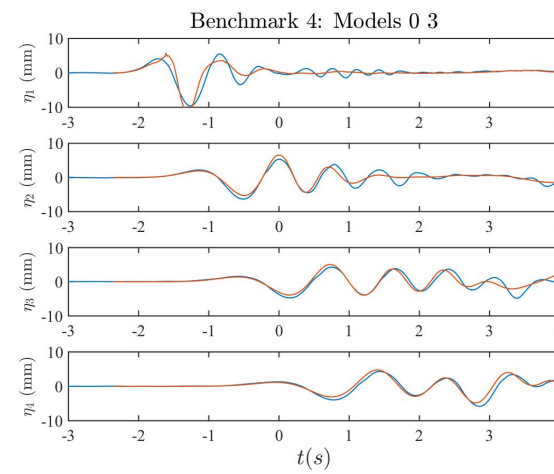
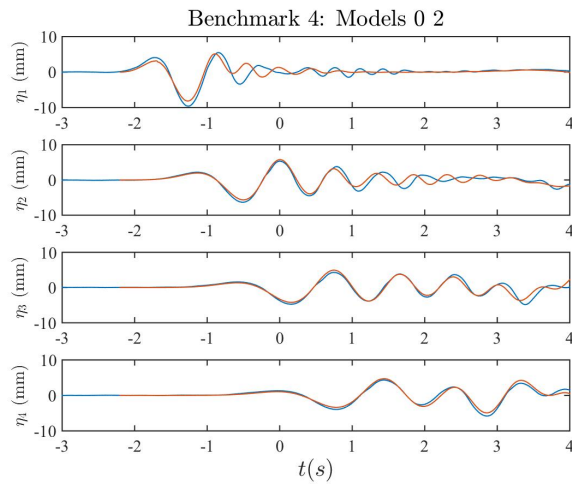


Benchmark 4: Models 0 16





Benchmark 4: Individual results for dispersive models





Error estimates

$$\epsilon_j = \frac{\sqrt{\sum_i (\eta_m(i, j) - \eta_o(i, j))^2}}{\max_i(\eta_o(i, j)) - \min_i(\eta_o(i, j))}; j = 1, 4$$

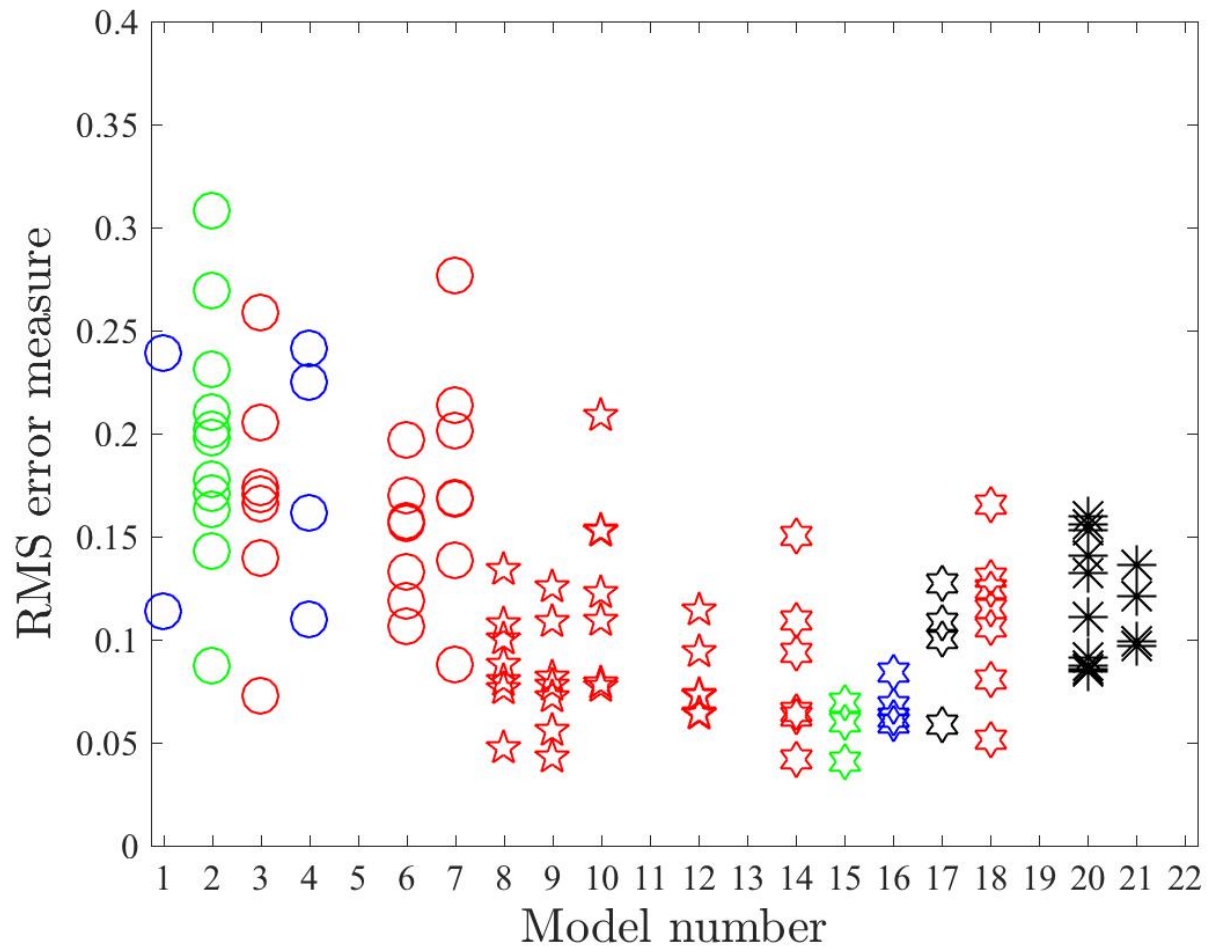
Subscript m denotes mode, o denotes observation.

Sum is over time series index i .

Individual gauges are denoted by j

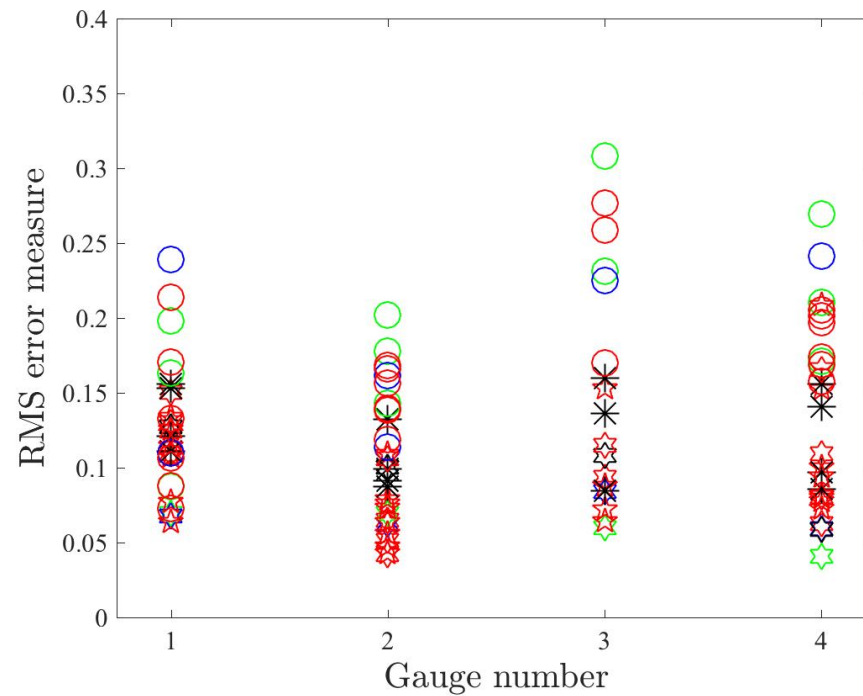


Distribution of errors by model for all gauges



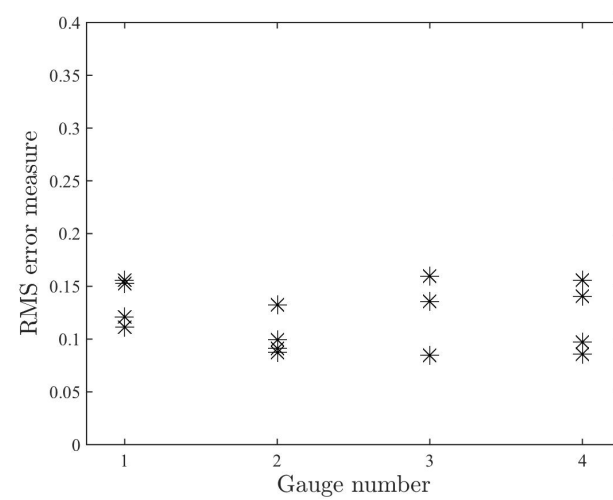
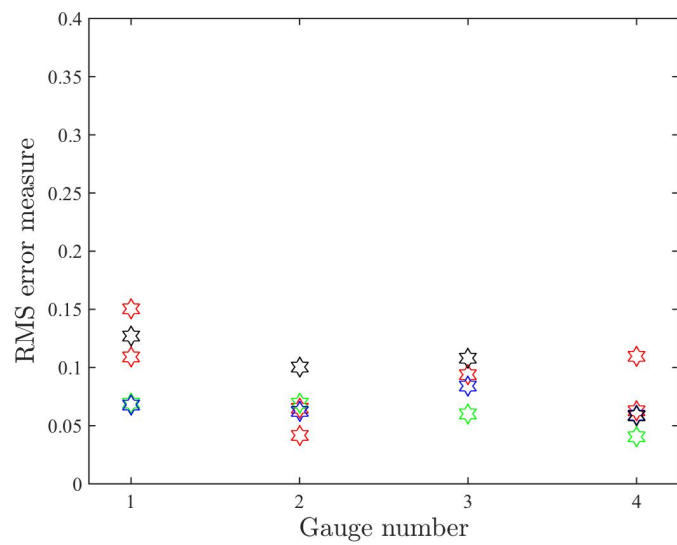
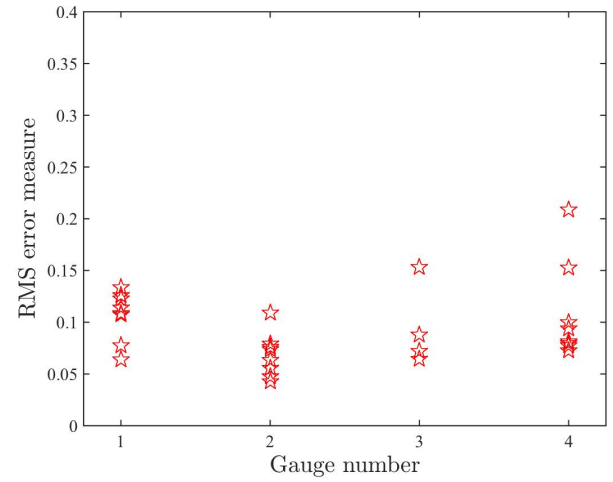
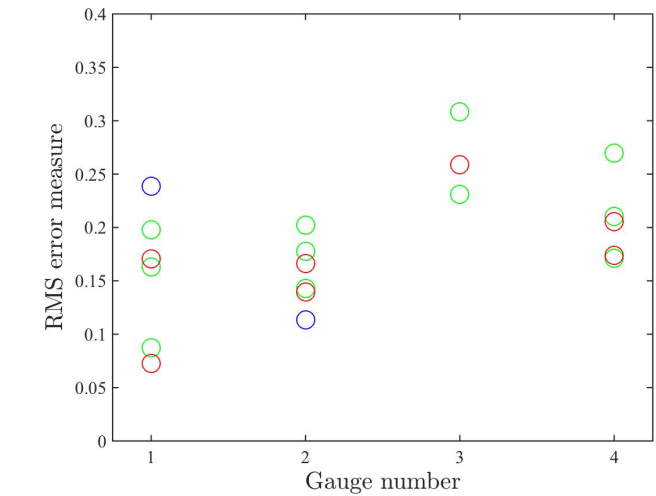


Distribution of errors for all models by gauge





By hydrodynamic type





Statistics of surface displacements

- Maximum surface displacement for a single gauge, normalized by observed values

$$\epsilon_1 = \text{Max}(\eta_{j,m}(t)) / \text{Max}(\eta_j^{obs}(t))$$

- Minimum surface displacement for a single gauge, normalized

$$\epsilon_2 = -\text{Min}(\eta_{j,m}(t)) / \text{Min}(\eta_j^{obs}(t))$$

- Total surface excursion for each gauge, normalized

$$\epsilon_3 = [\text{Max}(\eta_{j,m}(t)) - \text{Min}(\eta_{j,m}(t))] / [\text{Max}(\eta_j^{obs}(t)) - \text{Min}(\eta_j^{obs}(t))]$$

- Variance $^{1/2}$, normalized

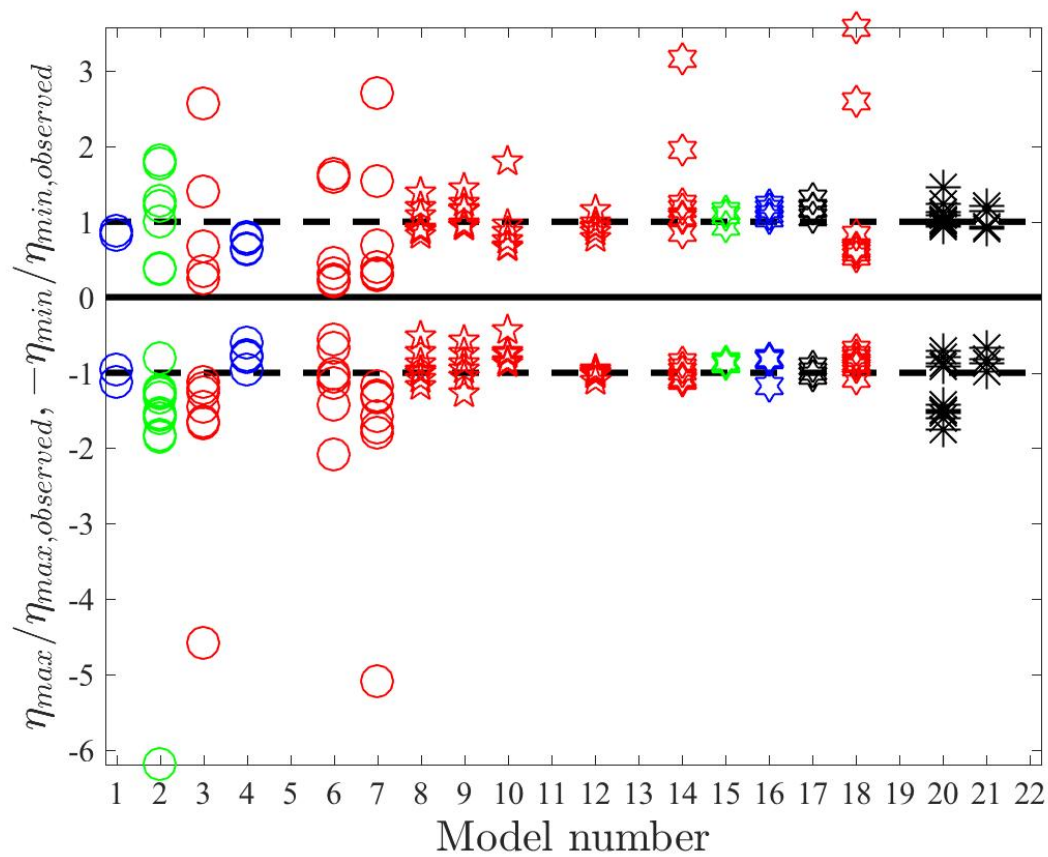
$$\epsilon_4 = \left[\int \eta_{jm}^2 dt / \int \eta_j^{obs,2} dt \right]^{1/2}$$

$$j = 1 - 4$$

$$m = \text{model}$$

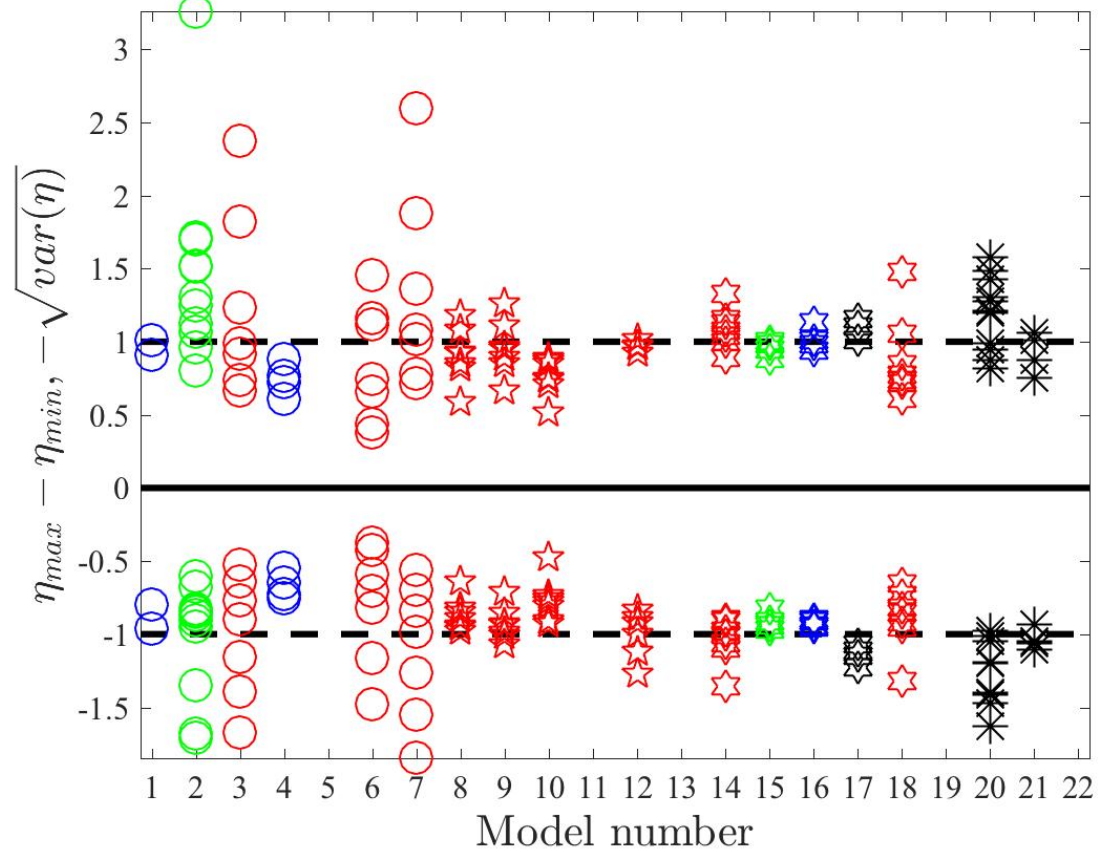


Normalized maximum positive and negative elevations



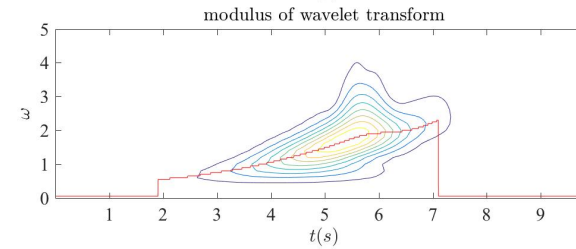
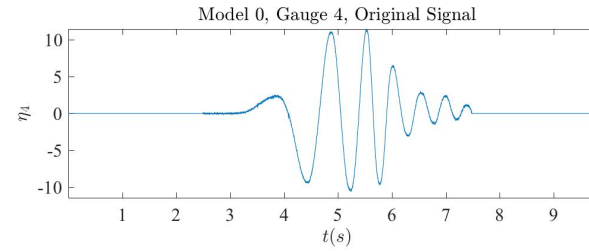
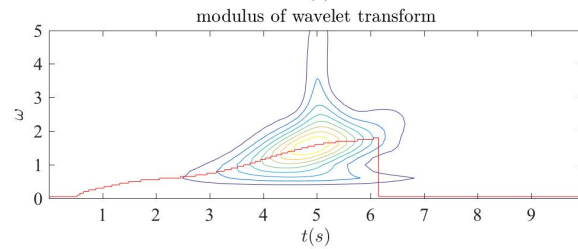
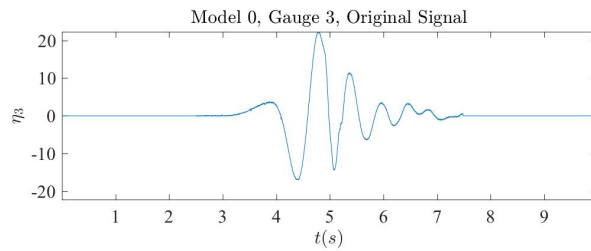
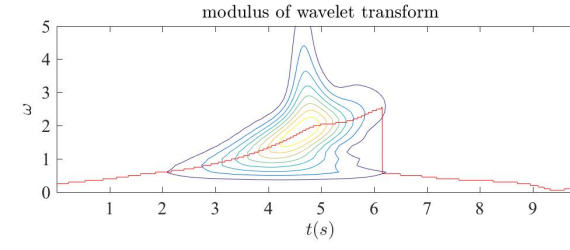
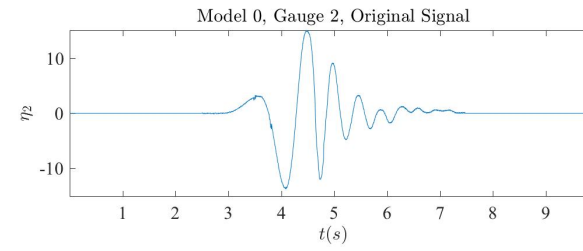
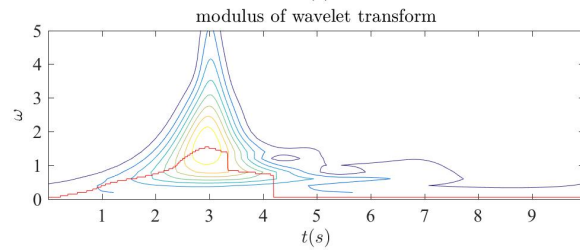
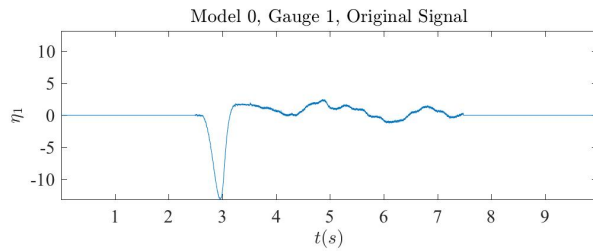


Normalized max excursion, variance^{1/2}





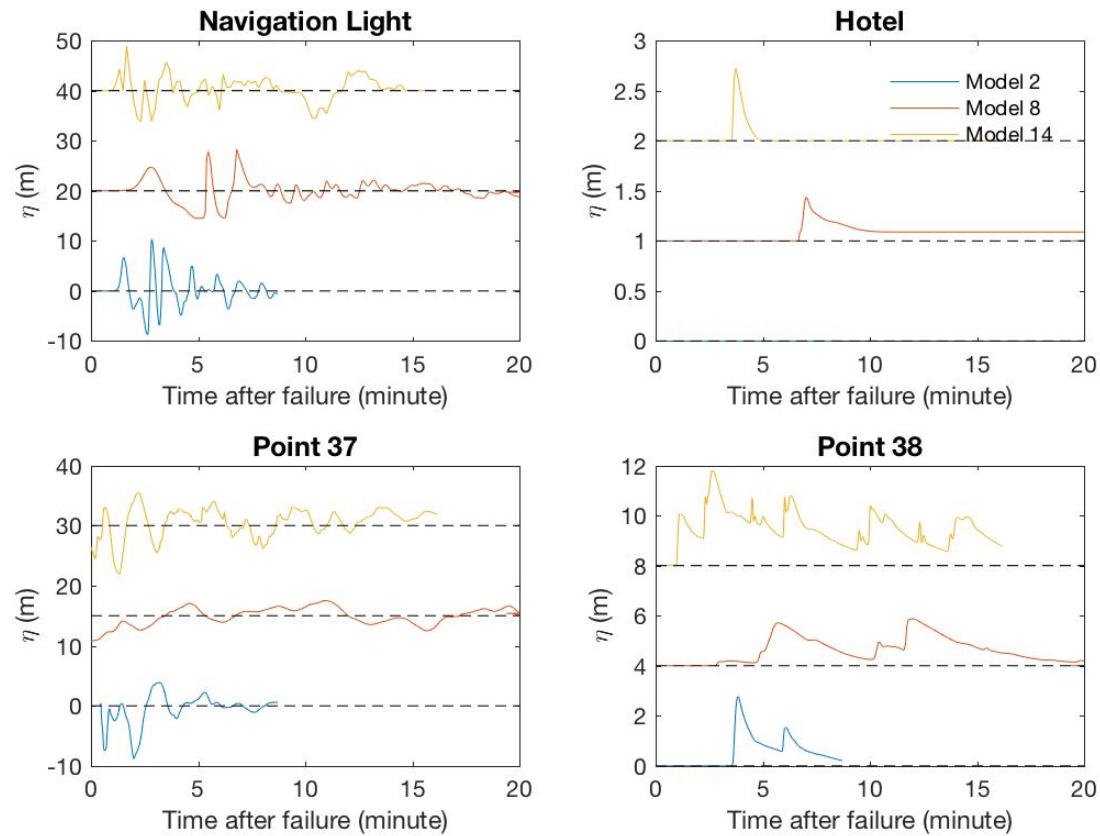
Structure of wave packet





Benchmark 7: Times series

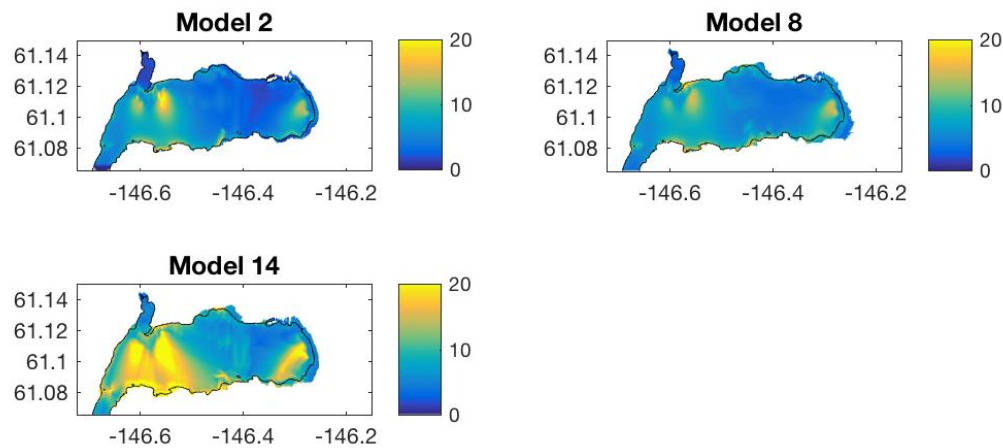
Benchmark 7: Models 2 8 14 22, Surface elevations





Variability of maximum surface elevations

Benchmark 7: Models 2 8 14 22, Maximum runup map





Variability of inundation line estimates

